



WAXI- West African Exploration Initiative
IXOA- L'Initiative d'Exploration Ouest Africaine

Lower crustal exhumation under transtension during the Palaeoproterozoic Eburnean Orogeny in Sefwi belt, SW Ghana

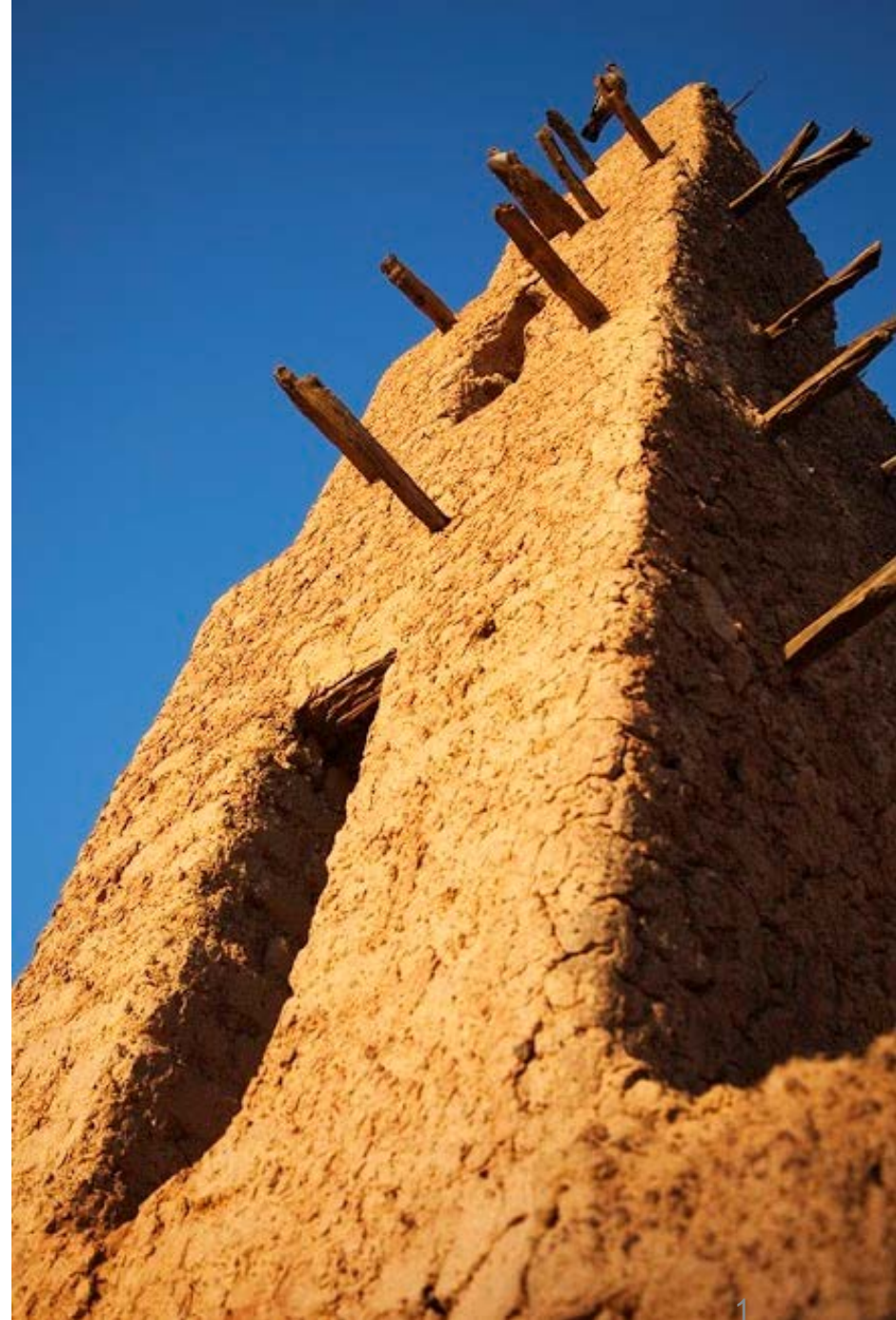
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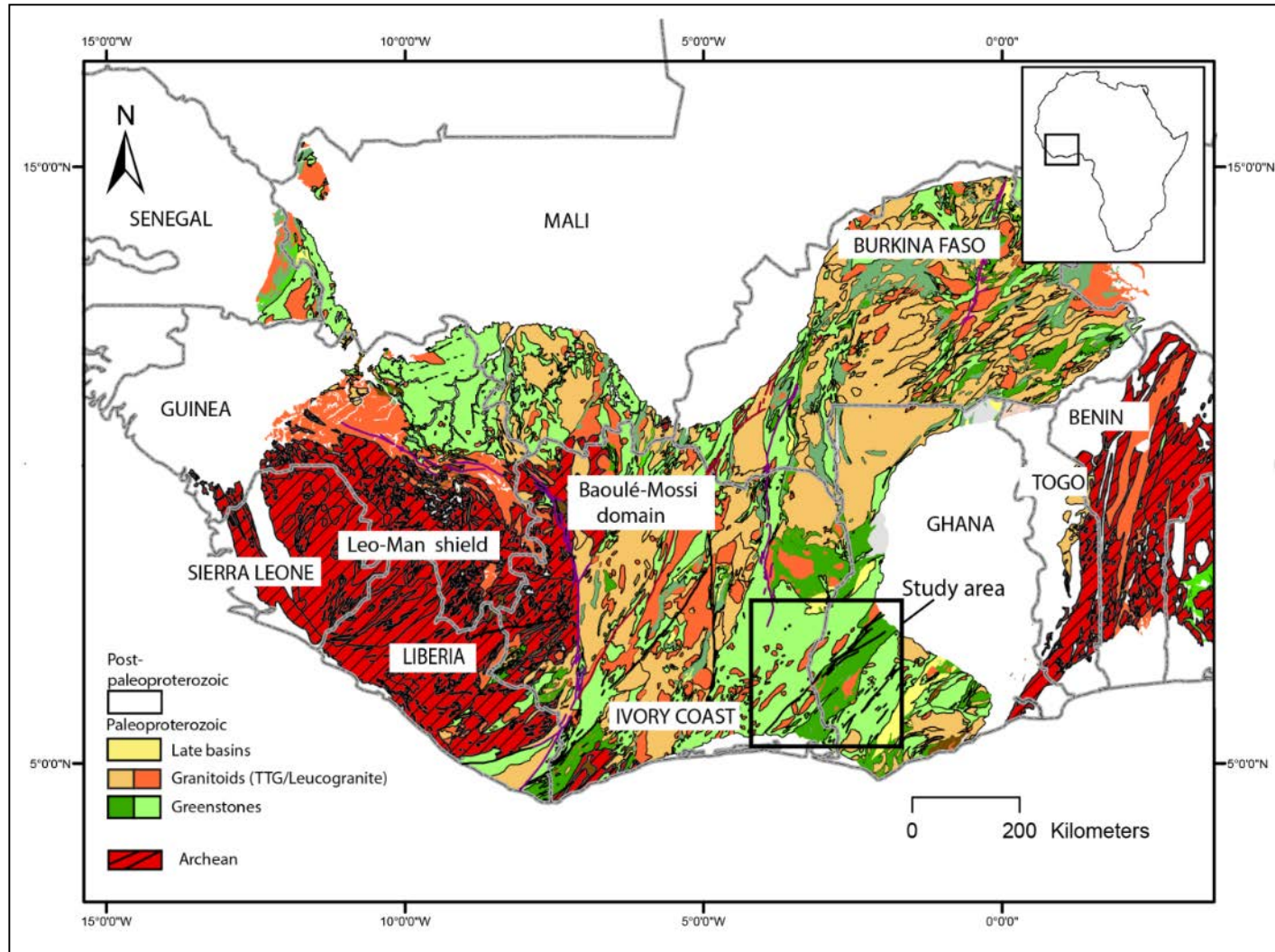
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Geological context



Paleoproterozoic Baoulé-Mossin domain in the southern portion of the West African Craton (modified after Milési et al. [2004](#), BRGM SIGAfrique).

Magmatic accretion of the juvenile crust of the WAC (~2.20 – 2.15Ga)

Varied models for the Eburnean Orogeny across the WAC

Monocyclic Eburnean Orogenic event

High grade rocks juxtaposed against low grade supracrustal equivalents by differential exhumation

(eg Bessoles 1977; Leube et al. 1990; Eisenlohr and Hirdes 1992; Hirdes et al. 1996; Hirdes et al. 2007; Block et al. 2015)

Multiple, discrete tectono-thermal events.

Eoeburnean (Eburnean I, Tangaeen) deformation (~2.20– 2.15 Ga) evident in high grade rocks followed by the Eburnean, separated by extension (~2.13 – 1.98 Ga)

(eg Lemoine et al., 1990; De Kock et al. 2011; 2012; Hein, 2010; Feybesse et al., 2006; Perrouty et al. 2012)

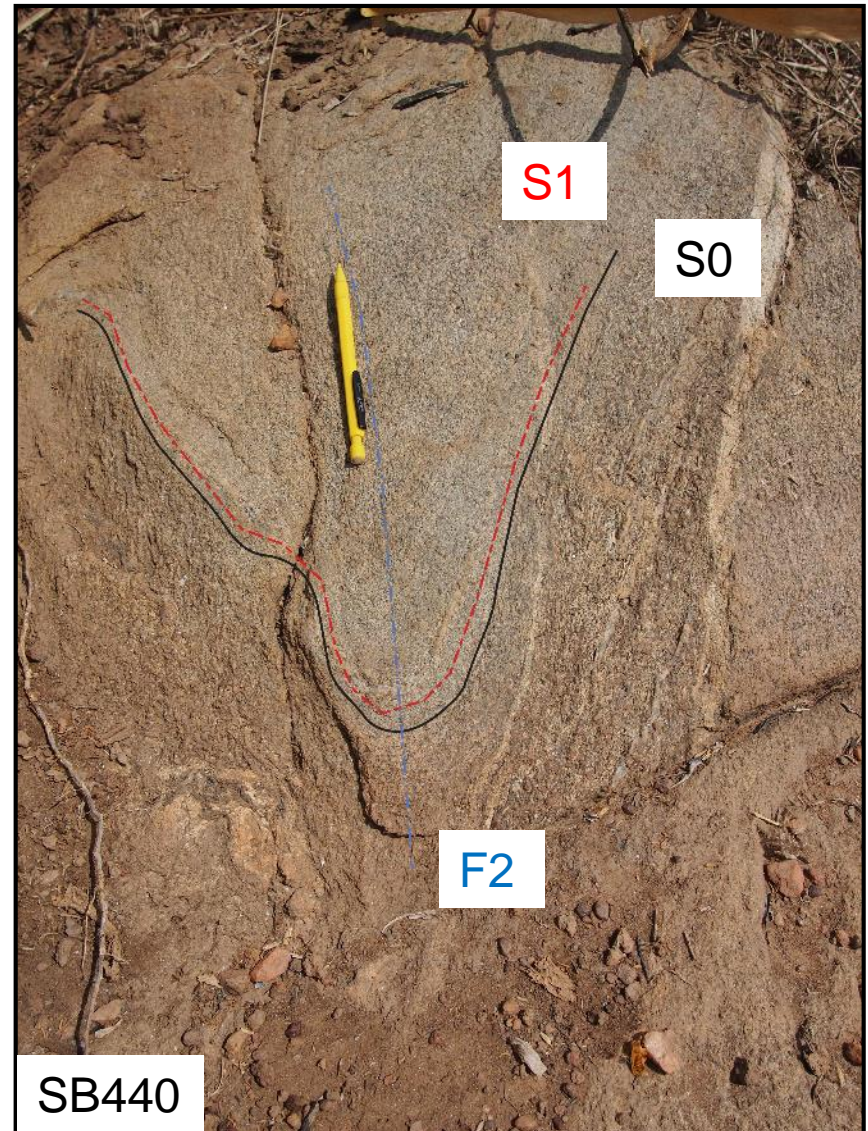
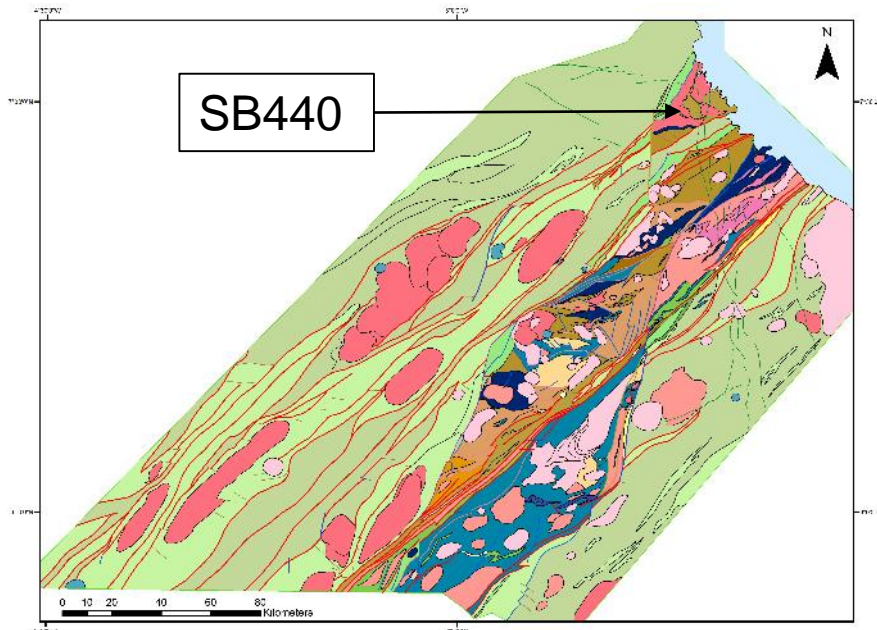
Subsequent transcurrent tectonism

Early N-S sinistral shear zones in C.I., Senegal and western B.F. overprinted by NE-SW dextral shear zones in western Ghana and eastern B.F.

(Ledru et al. 1991; Feybesse et al. 2006; Pouclet et al. 2006; Perrouty et al. 2012; Jessell et al. 2012)

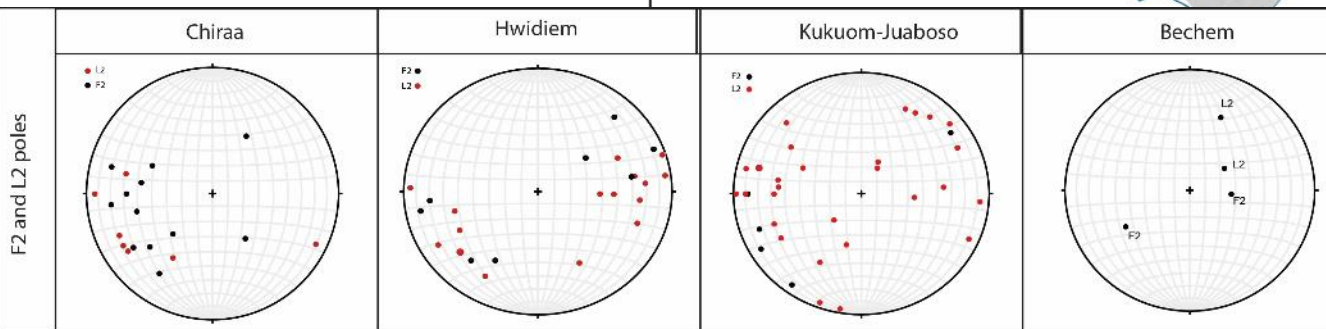
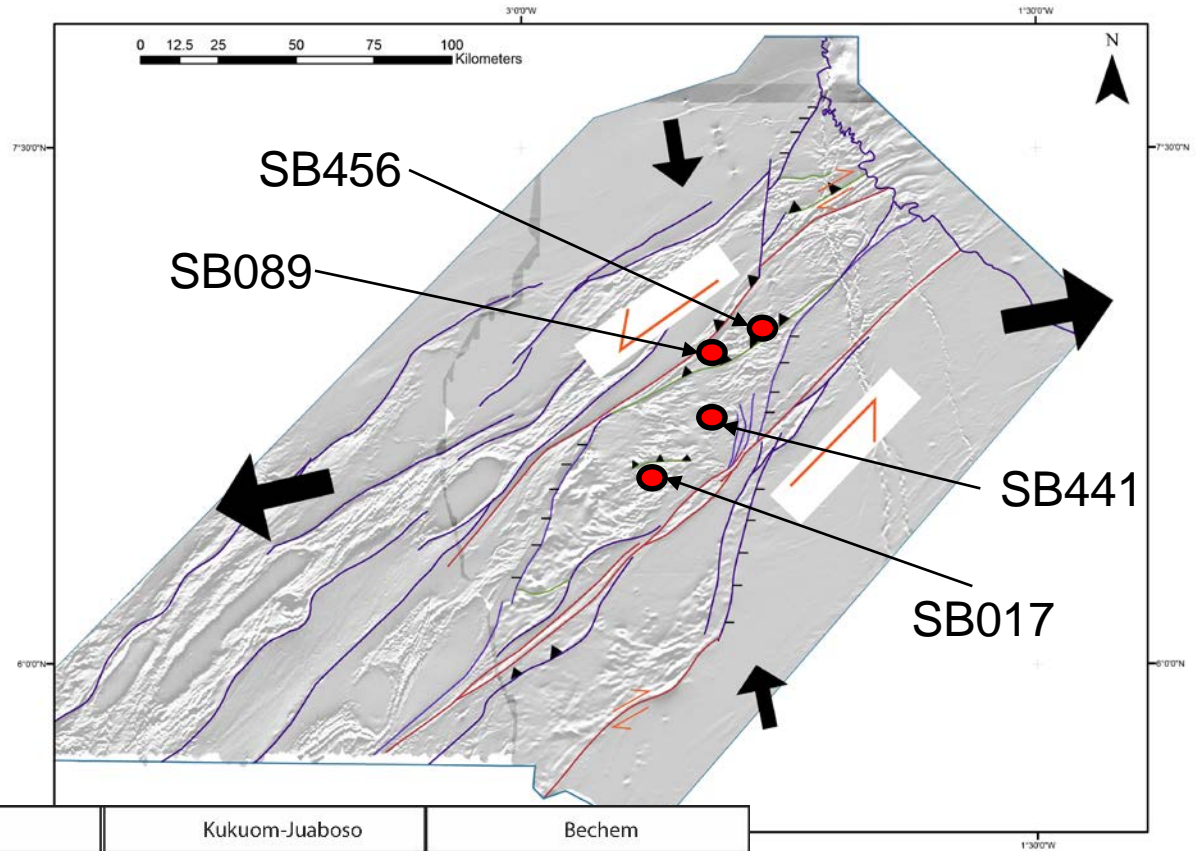
D1 Crustal thickening event

S_1 Bedding-parallel metamorphic foliation in schists, paragneisses



D2 transtension

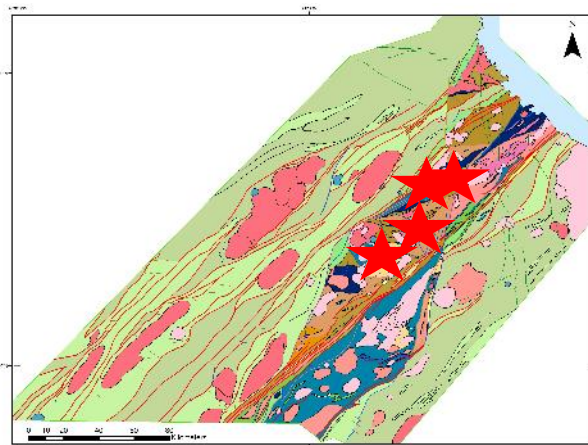
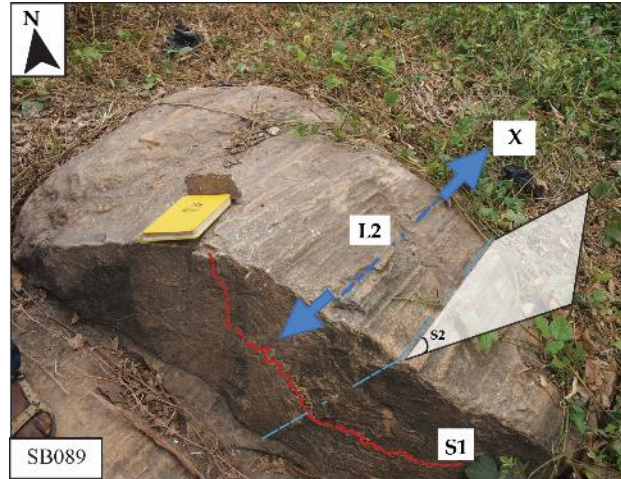
Structural domains of the belt show close relationship between L2 and F2



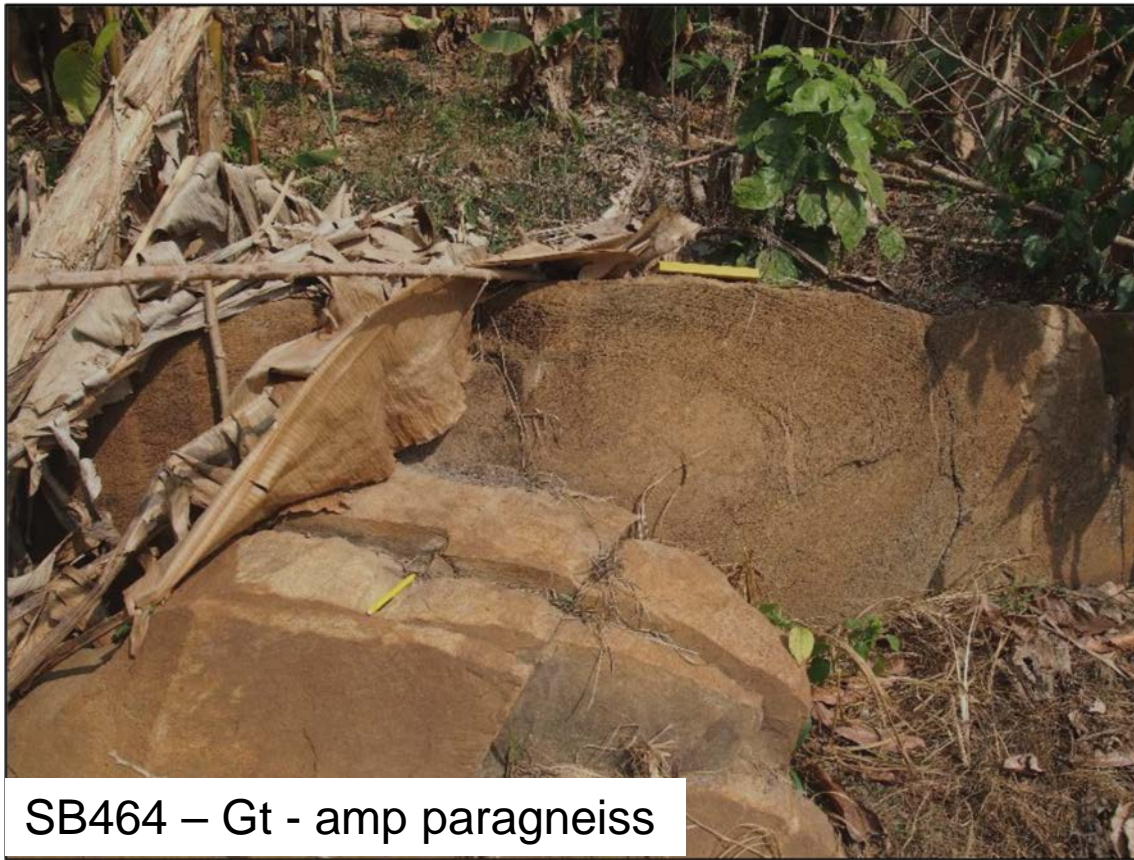
D2 ENE-WSW sinistral transtensional event

Open to tight E-W to
NE-SW
F2 folds

F2 approx. = L2
mineral stretching
lineation



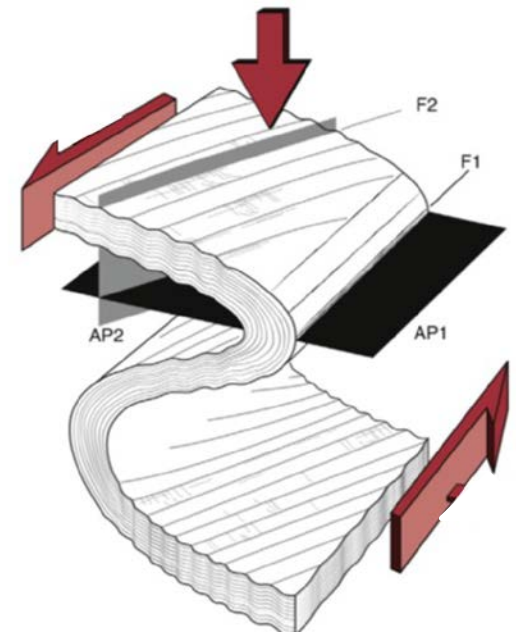
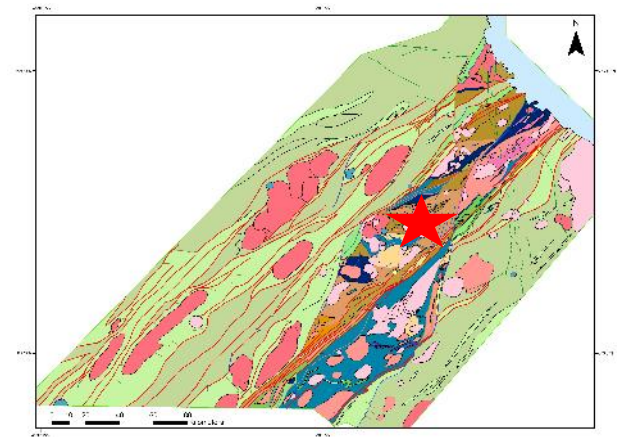
Further evidence of transtension



SB464 – Gt - amp paragneiss

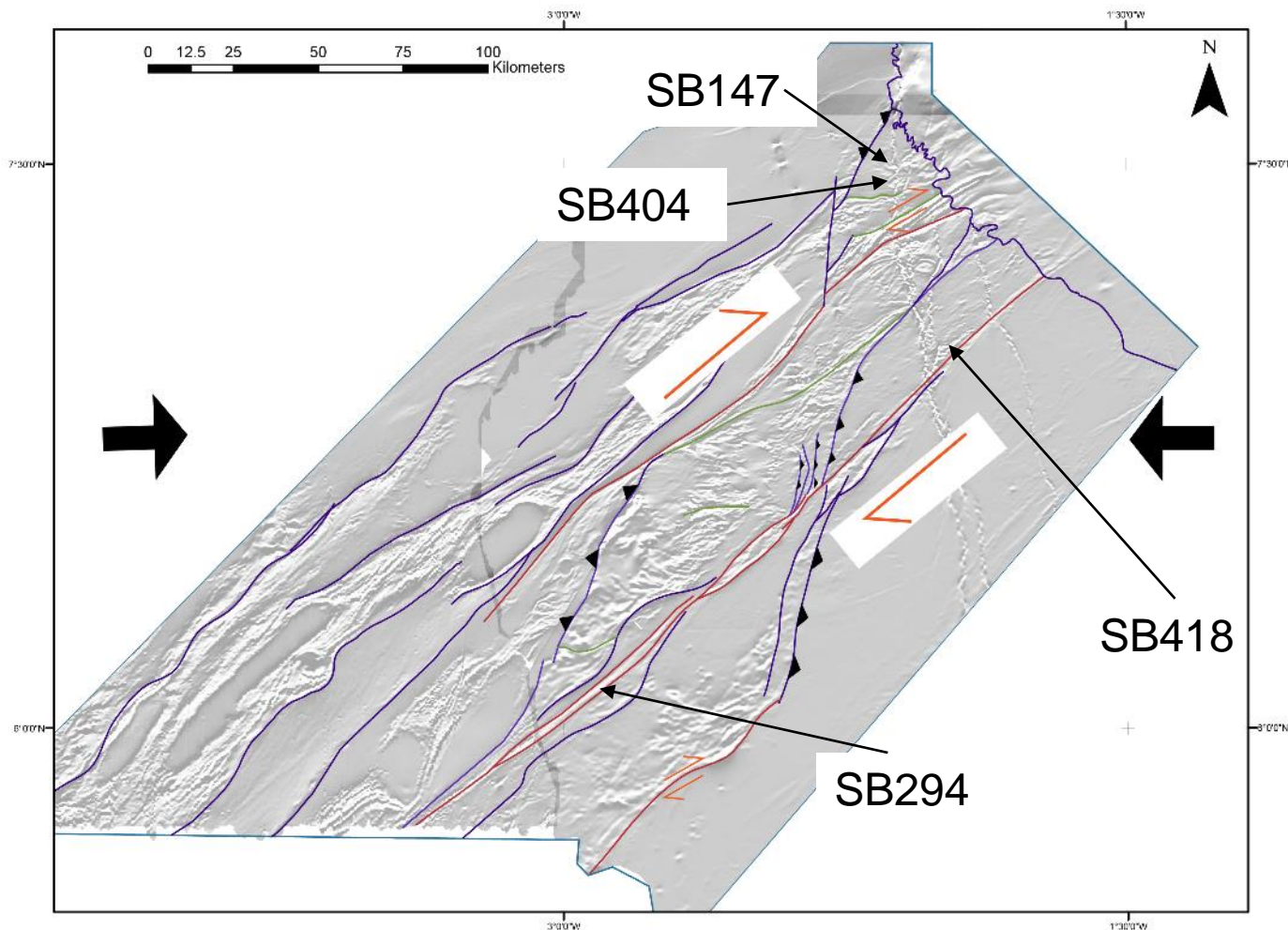
Fold superposition during progressive transtension.
NB Pencil on subhorizontal L2 intersection lineation
(10 → 250)

F2 10 → 230



From Fossen et al 2013

D3 – Refolding of F2s and reactivation of pre-existing D2 faults as thrusts



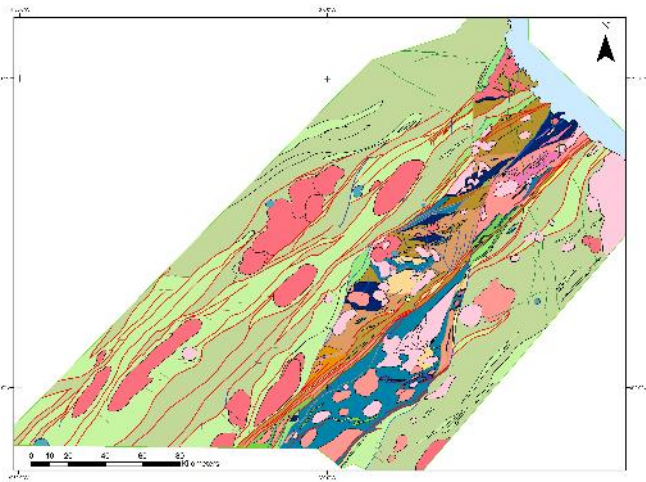
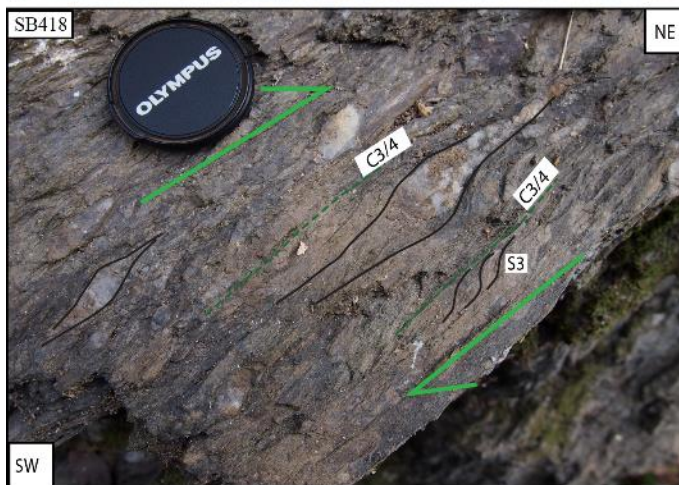
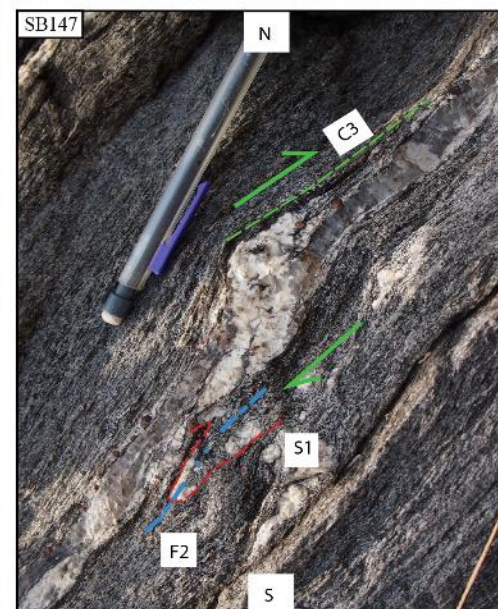
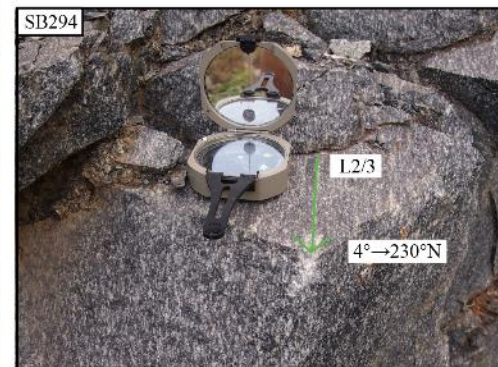
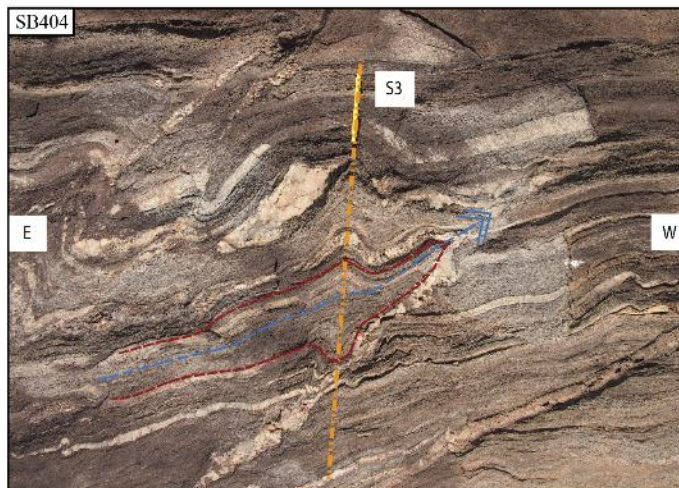
Late D3/4 continued strain localization along narrow high strain zones

Structural observation: D₃ / 4

N-S upright F₃ folds
overprint E-W folds

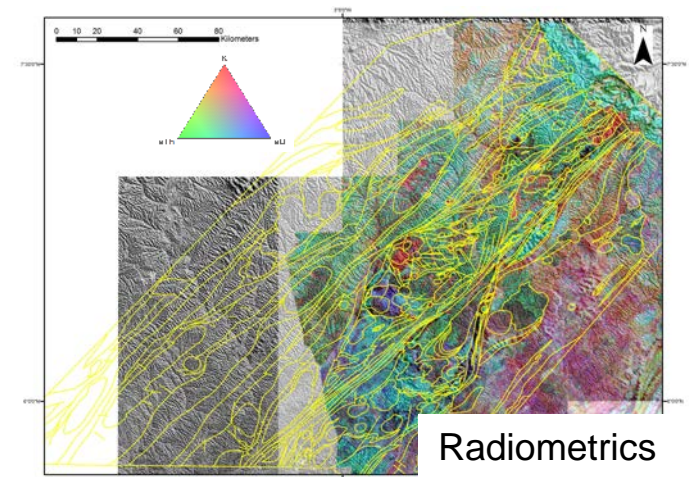
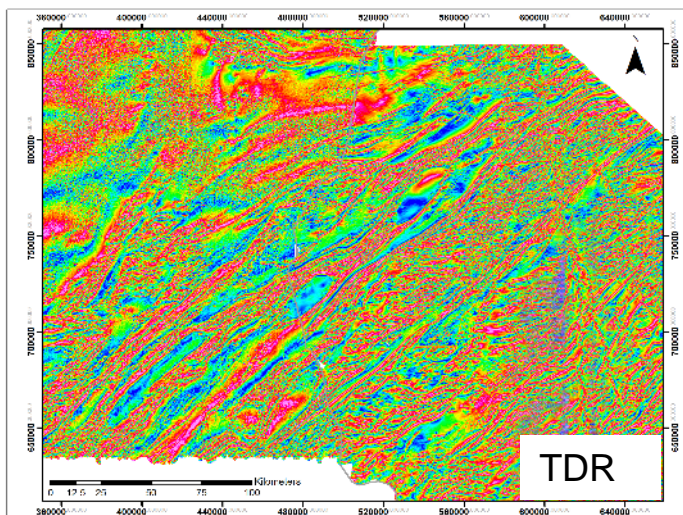
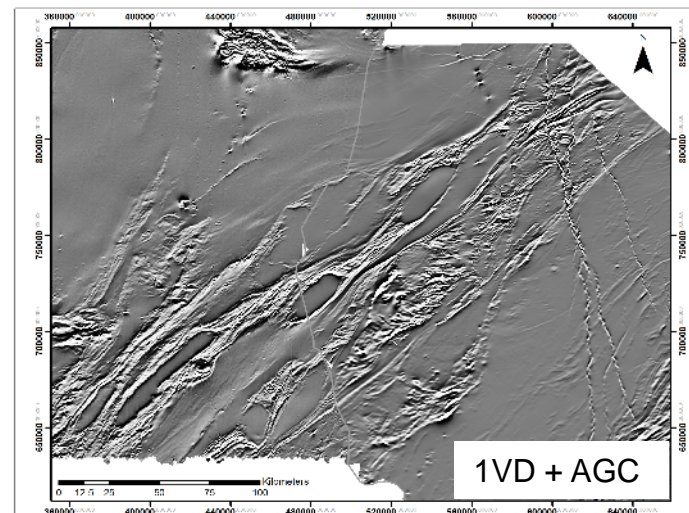
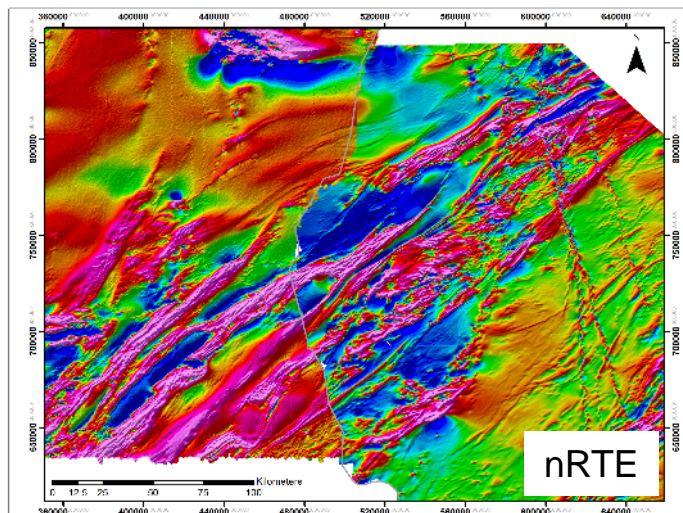
S₃ defined by N-
striking crenulation
cleavage
and NE-SW striking
S-C_{3/4} fabric in
shear zones

L-tectonites and S-
C₃ fabrics



Regional geophysical datasets used to interpret the regional scale structures

- 484 outcrop localities
- Newmont and Kinross outcrop databases
- Measurements of the petrophysical properties of the major lithologies



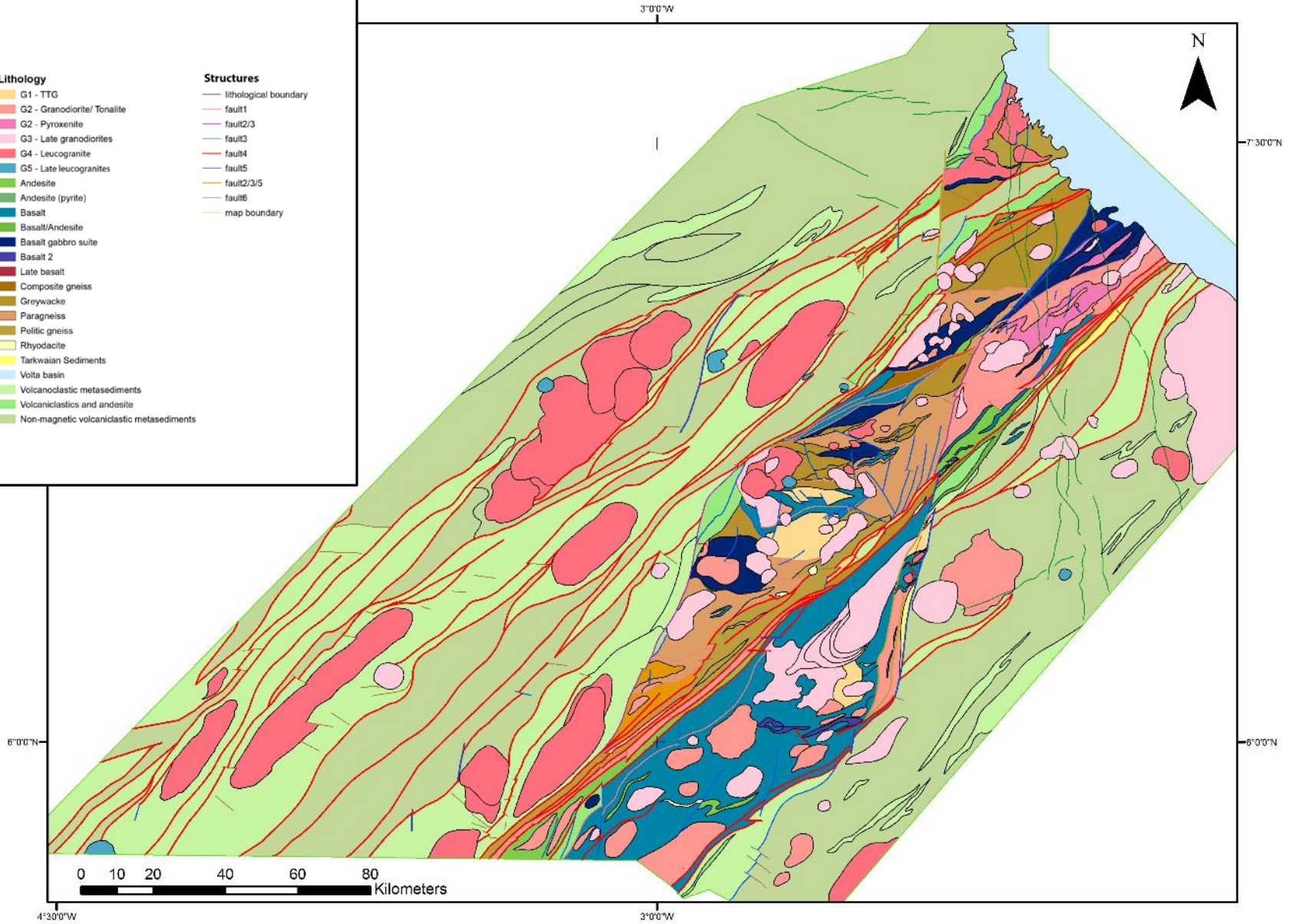
Litho-structural map

Lithology

- G1 - TTG
- G2 - Granodiorite/ Tonalite
- G2 - Pyroxenite
- G3 - Late granodiorites
- G4 - Leucogranite
- G5 - Late leucogranites
- Andesite
- Andesite (pyrite)
- Basalt
- Basalt/Andesite
- Basalt gabbro suite
- Basalt 2
- Late basalt
- Composite gneiss
- Greywacke
- Paragneiss
- Pelitic gneiss
- Rhyodacite
- Tarkwaian Sediments
- Volta basin
- Volcanoclastic metasediments
- Volcaniclastics and andesite
- Non-magnetic volcaniclastic metasediments

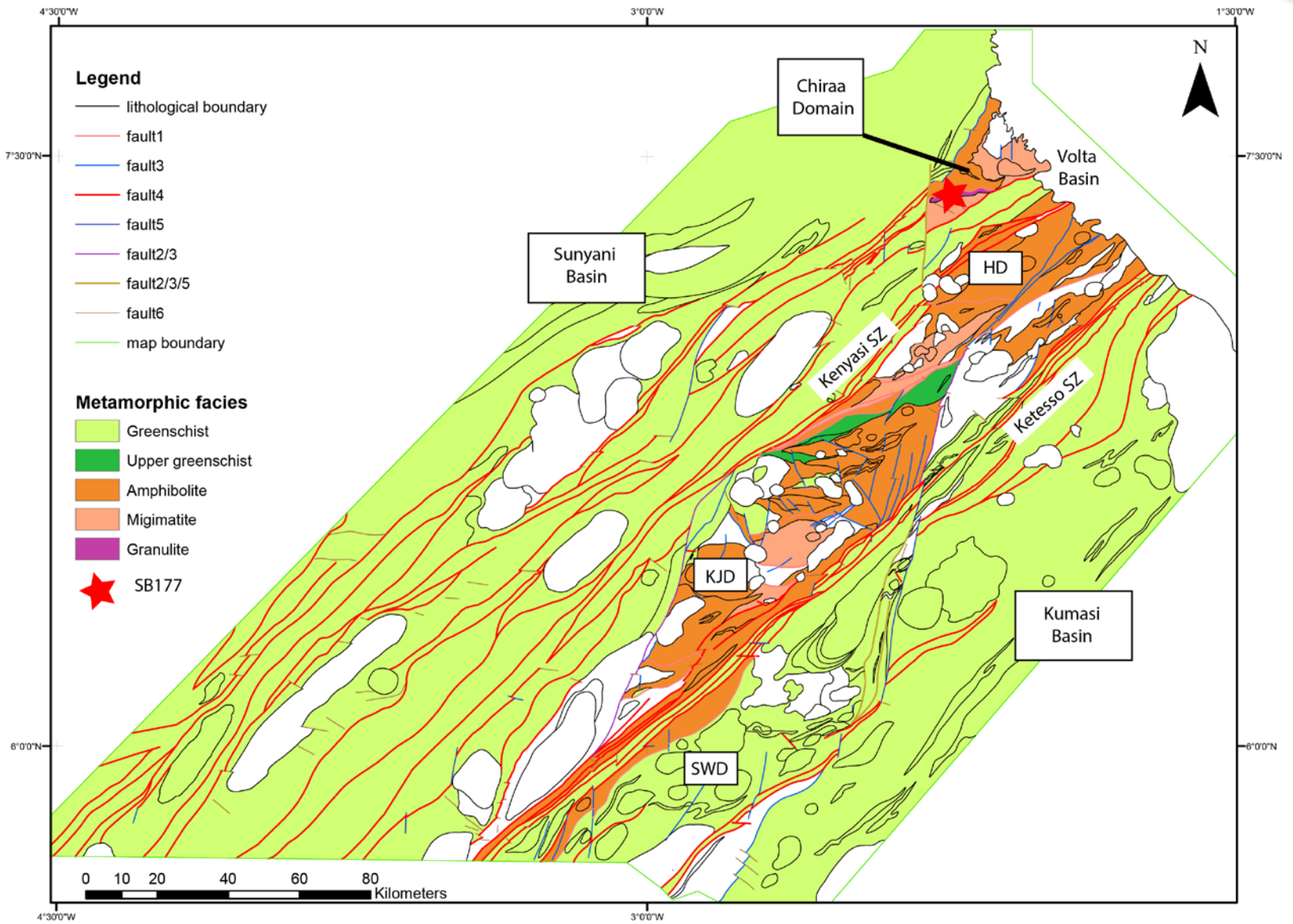
Structures

- lithological boundary
- fault1
- fault2/3
- fault3
- fault4
- fault5
- fault2/3/5
- fault6
- map boundary



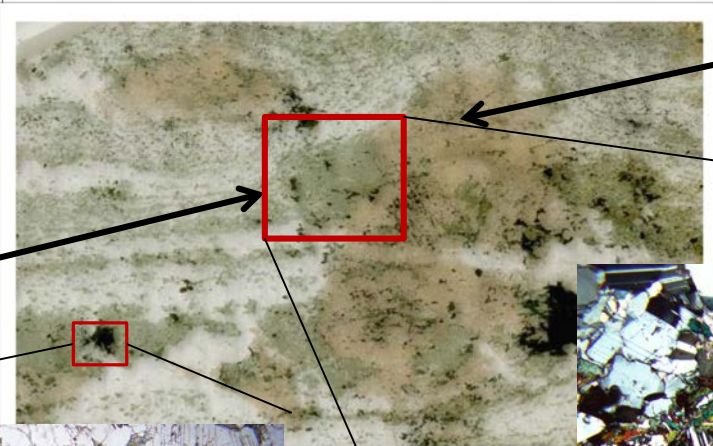
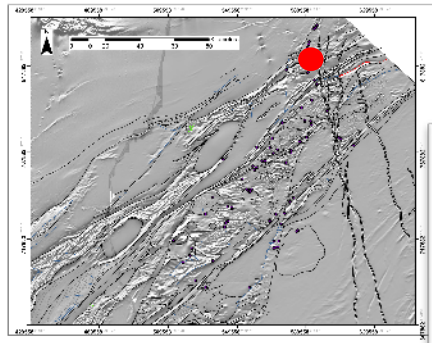
- D1 Early gneissic or layer parallel foliation
 - Burial of supracrustal rocks during crustal thickening
- Anatexis: Late D1 to syn-D2
- D2 ENE-WSW transtension (with NNW-SSE shortening event)
 - Tight, upright to inclined E-W to ENE-WSW F2 folds and oblique thrust faults
 - Prominent elongation along F2 fold axes (L2 subparallel with F2)
 - Oblique sinistral reverse movement on major shear zones
 - Amphibolite facies conditions
- D3/D4 E-W shortening event (D3)
 - Regional, narrow, NE-SW high strain dextral shear zones
 - Local NE-SW S4-C fabrics, including transposition early structures
 - GS facies metamorphic overprint
 - Reactivation of D2 faults as reverse faults
- D5 Late NNE-SSW strike slip faults with apparent sinistral displacement

Metamorphic map



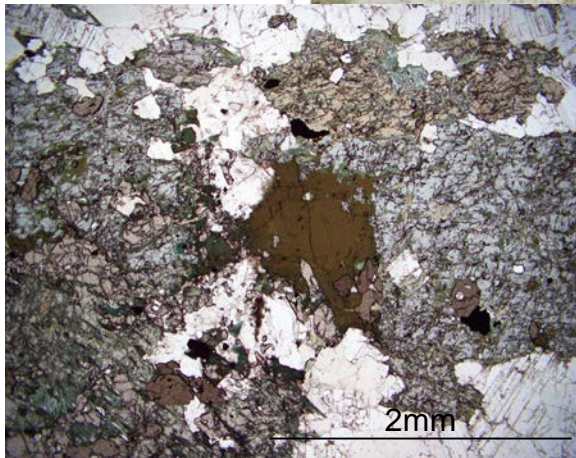
SB177: Mafic Granulite

Peak assemblage: Cpx + Pl + Gt + Spn(Tnt) ± Qtz + melt

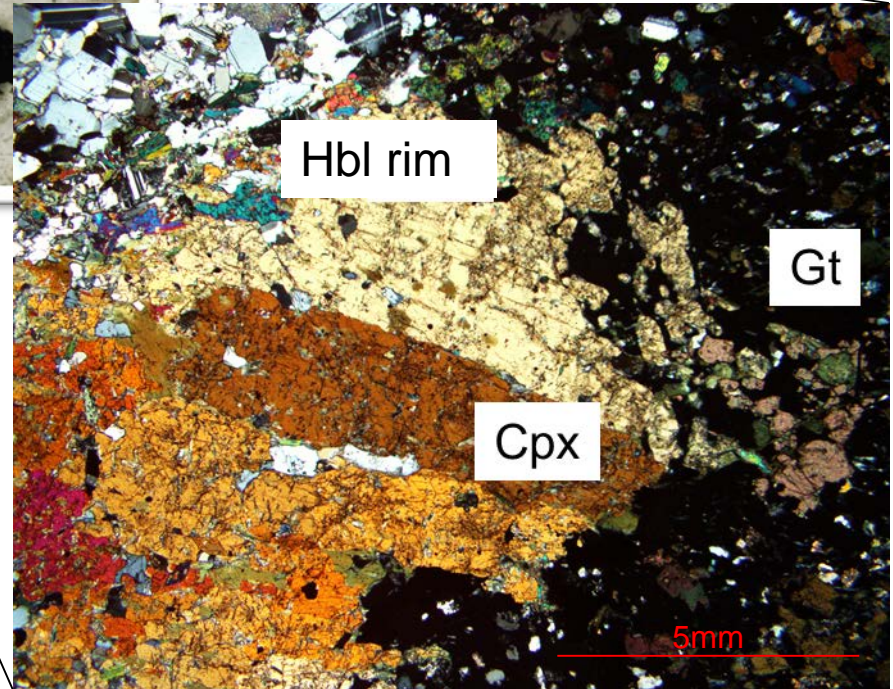


Gt porphyroblast

Cpx



Hbl in boudin neck of Cpx
= Retrogression



Hbl rim

Gt

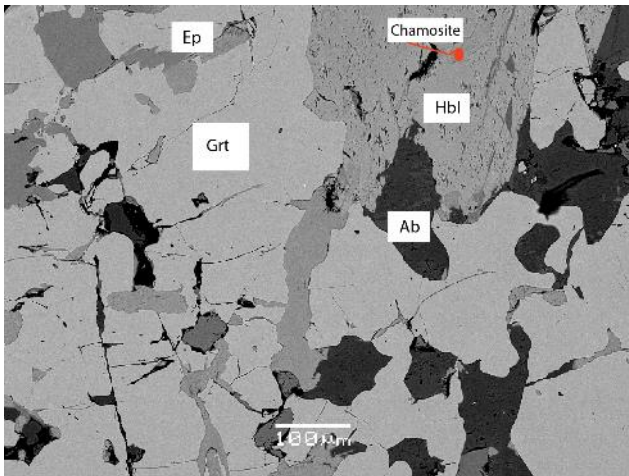
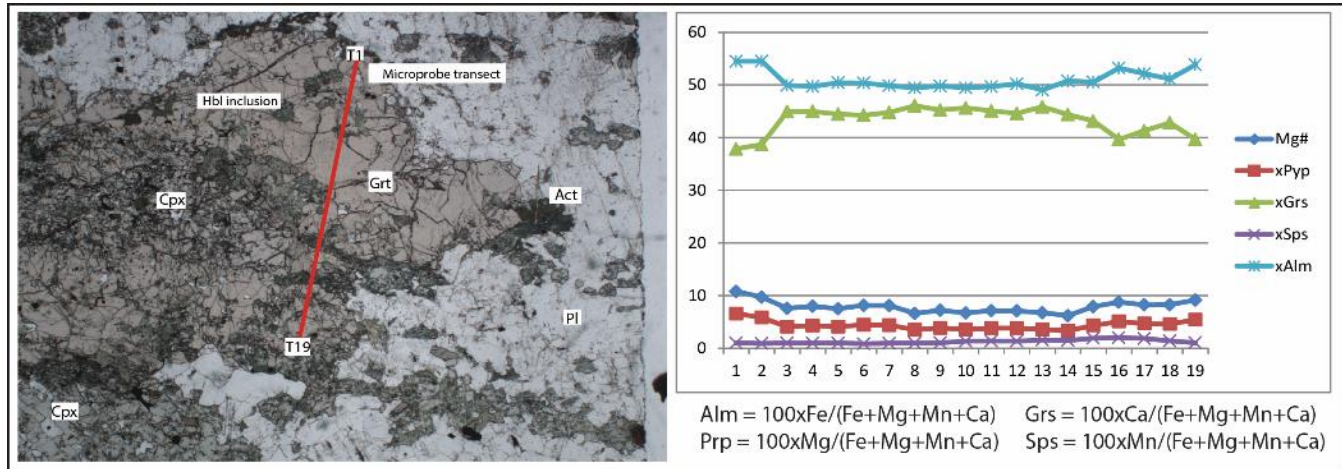
Cpx

5mm

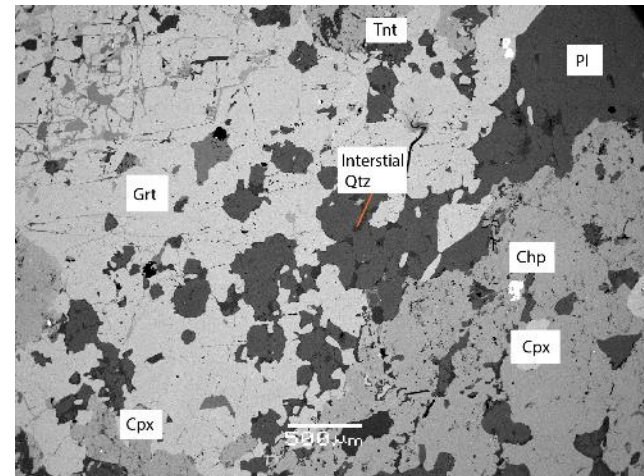
High T shearing: Cpx in strain shadow of garnet porphyroblast

Pseudosection constraints

Radial transect of Grt porphyroblast showing chemical composition variation



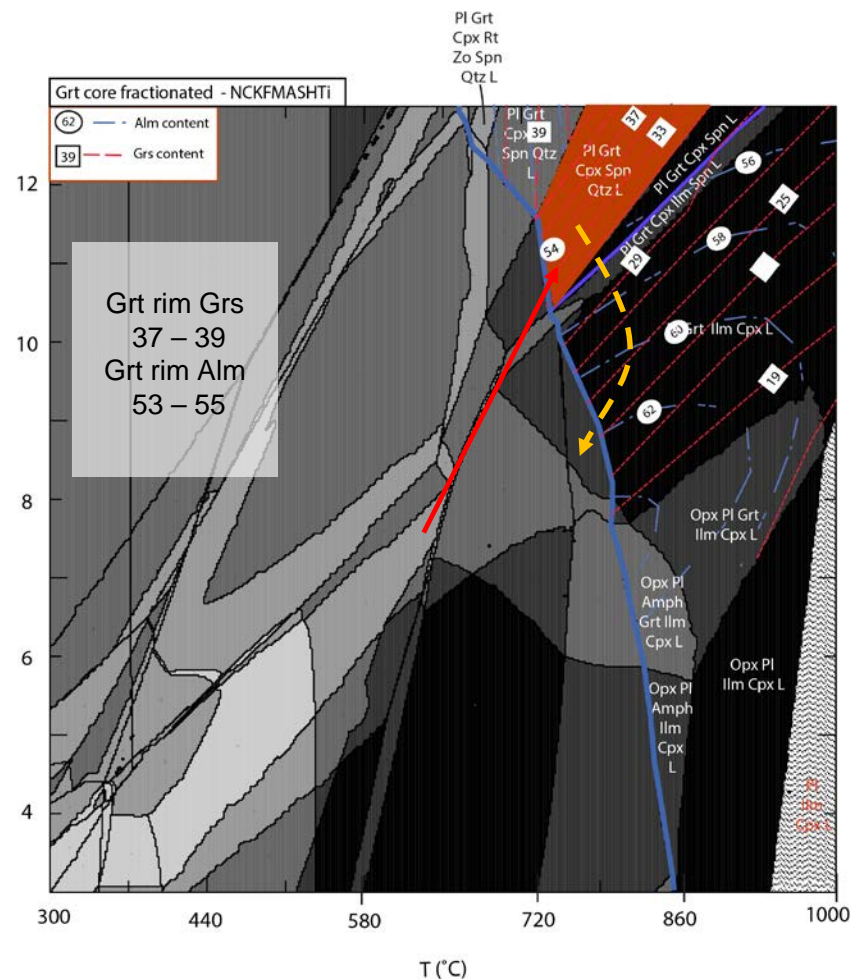
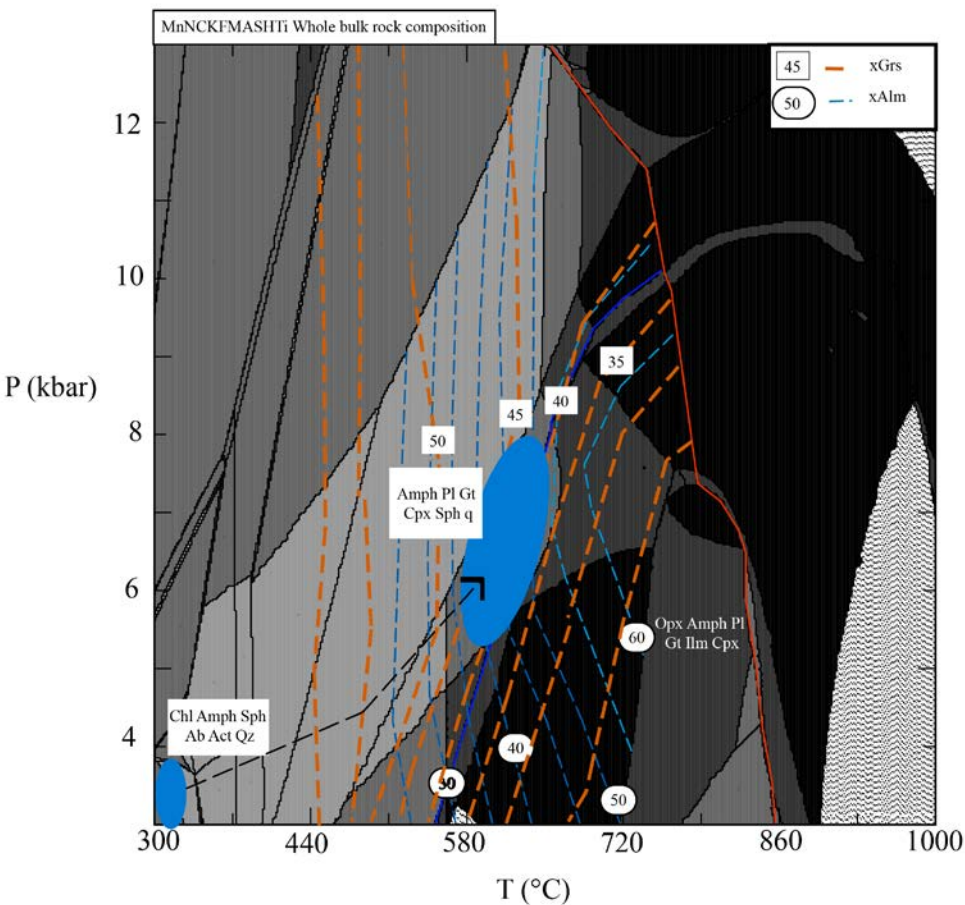
a) BSE image showing M1e greenschist to amphibolite facies assemblage preserved in Grt core inclusions



b) BSE image of rim of Grt porphyroblast intergrown with Pl and in contact with Cpx porphyroblast in leucosome (peak assemblage)

a) Prograde path

b) Peak P-T after grt core fractionation



For mafic granulites at $P > 10\text{kbar}$, reaction (R1) $\text{Hbl} + \text{Pl} + \text{Qtz} = \text{Grt} + \text{Cpx} + \text{L}$
(Pattison et al. 2003)

- Peak metamorphic conditions of buried supracrustal rocks in belt record anatexis at 10 – 12 kbar, 700 - 800°C → associated with major crustal thickening D1
- Exhumation of high grade rocks eg. mafic granulite during D2 transtension
- Three major, discrete tectonic events produced the regional architecture of the Sefwi belt.
 - D1 - crustal thickening driven by external tectonic forces
 - D2 - transtensional event
 - D3 - transpressional event
- Tectonic juxtaposition of high grade rocks and greenschist facies lithologies prior to regional shearing (D4)
- First evidence of exhumation by transtension proposed for WAC during the Palaeoproterozoic

- Results show contrasting strain regimes, metamorphic histories and exhumation methods records in WAC greenstone belts
 - Significant D1 crustal thickening evident in both Sefwi (this study) and Wa-Bole regions (eg. Block et al., 2015)
 - D1/2 early fabric recognised in Sefwi Group in Ashanti belt (Allibone et al, 2002; Perrouty et al. 2012) and in this study
 - New exhumation mechanism proposed for high grade rocks in Sefwi belt – transtension (extension axis $\sim 90^\circ$ to that in NW Ghana)
 - Ashanti belt shows prominent NW-SE shortening in D3 and subsequent sinistral reactivation of the Ashanti fault (Perrouty et al. 2012), where D3 is represented by long lived E-W shortening, however timing unconstrained thus far.

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MONASH
University



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